

SUBSURFACE GEOLOGY OF THE ST. CROIX CARBONATE SYSTEM  
PHASE II

Ivan P. Gill  
Dennis K. Hubbard

May, 1987

Agreement No. 14-08-0001-G1258

Technical Report No. 28  
Caribbean Research Institute  
University of the Virgin Islands  
St. Thomas, U.S.V.I. 00802

Technical Report No. MG-4  
West Indies Laboratory  
Teague Bay, St. Croix  
U.S. Virgin Islands 00820

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The research on which this report is based was financed in part by the United States Department of the Interior, Geological Survey, through the Virgin Islands Water Resources Research Center.

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## ABSTRACT

Seven new test holes were drilled into St. Croix's central carbonate plain with a rotary drill rig. Cumulative drilling for this phase of the project exceeded 750 feet, and brought the total number of test holes for the project to fourteen. These test holes allow lithologic and biostratigraphic correlation in a north to south transect from Krausses Lagoon to Estate St. John and in a west to east transect from Estate Hesselberg to Estate Pearl.

The drilling establishes the existence of a probable Pliocene reef and shallow-water facies trend that rims the western and southern coastlines of the central plain. The greatest thickness of Pliocene sediments occurs in a subsidiary graben block in the south coast industrial area. The northern and western boundaries of the Pliocene graben can be inferred from core data. The Pliocene post-Kingshill carbonates are less extensive than the Miocene Kingshill Limestone, but are generally more permeable.

Dolomitization in the Pliocene carbonates rims what was the coastline of Krausses Lagoon before industrial development modified the shoreline in the 1960s. The geographical distribution of the dolomite suggests a hydrologic correlation between Krausses Lagoon and the formation of dolomite. The stable isotopic composition of the dolomite suggests the possibility of a dolomitizing fluid with elevated salinity.

Structural mapping on the upper surface of the Miocene Jealousy Formation indicates marked upwarping under the carbonate highlands. This structural upwarping coincides with the greatest isopach thickness of the Kingshill Limestone. The patterns suggest a basin opening to the south, but with a depocenter located under the present position of the carbonate highlands. There is a greater degree of structural complexity in the central plains region than was previously supposed.

Micropaleontological evidence suggests that the Jealousy Formation / Kingshill Limestone contact is time-transgressive within the Miocene. Both units were deposited in deep water, perhaps at depths greater than 1000 m. Despite the marked color change, there are surprisingly few mineralogic or paleontological differences between the Jealousy Formation and the Kingshill Limestone, and the contact between the two formations does not imply significant paleobathymetric change.

## ACKNOWLEDGEMENTS

This project has been made possible through the efforts of many people and the cooperation of numerous agencies. Funding was provided by the United States Department of the Interior through the Virgin Islands Water Resources Research Center; SOHIO, Chevron, and Shell field research grants; grants from Dr. David Eby and Champlin Petroleum, and the Applied Carbonate Research Program, the Department of Geology and the Basin Research Institute at Louisiana State University. Initial field work was funded by grants from the Geological Society of America and the American Association of Petroleum Geologists. The drilling would not have been possible without the aid and cooperation of Mr. Ken Eastman and the staff of Caribbean Drilling Services.

Access to exposures and drill sites for this phase of the project was freely given by the staff of Martin Marietta Corporation, in particular G. Bennewith and J. Savage, as well as H. Kerr, O. Schjang and the Women's Coalition of St. Croix. Cooperation during this phase of work was extended by several agencies of the Virgin Islands and Federal Governments: the Department of Public Works, the V. I. Planning Office, the Department of Natural Resources, and Mr. H. Rodrigues and the manbucket crew of the V. I. Water and Power Authority (St. Croix).

Field work for this phase of the project was aided generously by Y. Bordeaux, A. Hunt and J. Massare. Report preparation was aided by E. Babin, A. Brunett, and C. Van de burgh. Strontium isotopic work was generously donated by R. Koepnick and the Mobil Field Research Laboratory. Geophysical logging gear was loaned by Argonne National Laboratory courtesy of Mr. R. Bowen and Dr. L. McGinnis. Special thanks are owed to K. Carter, N. Martinez, K. Myers and M. Price for sample preparation and micropaleontological work, and to Sam Reed and T. Poche for thin section preparation. Micropaleontological determinations were done by P. McLaughlin and W. van den Bold, and the authors benefitted from discussions with R. Ferrell, R. Koepnick, E. Manning, P. McLaughlin, C. Moore, D. Nummedal, R. Pilger, B. Sen Gupta, M. Simms, W. van den Bold and S. Wendtler. S. Frost contributed enthusiasm, samples and ideas regarding the carbonate section of St. Croix.

We appreciate the support and advice of the staff of the U. S. Geological Survey, Puerto Rico, in particular Mr. F. Gomez-Gomez and Mr. A. Zack. Logistical support and management was provided by the staff of the West Indies Laboratory, and the program was administered by Dr. H. Smith of the Water Resources Research Center of the University of the Virgin Islands. One of the authors, Gill, is supported on a fellowship from the Louisiana State University Alumni Federation and the Department of Geology, and his lab work is supported by Dr. C. H. Moore and the staffs of the Applied Carbonate Research Program, the Basin Research Institute and the Department of Geology of Louisiana State University. The staffs of the Applied Carbonate Research Program and Department of Geology of Louisiana State University, and the staff of the West Indies Laboratory were invaluable in providing both field and laboratory assistance throughout the project.

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